## **Transition of CO<sub>2</sub> Enhanced Oil Recovery to Carbon Storage: Experimentally Constrained Reactive Transport Model**

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#### Schlumberger, "Carbonate Reservoirs," 2007.

## **Goals and Benefits**

- To quantify key relationships in reactive transport models to constrain final CO<sub>2</sub> storage estimates.
- To calibrate down hole logging measurement methods to estimate carbonate formation permeability.
- Our results improve prediction of changing CO<sub>2</sub> storage capacity in carbonate reservoirs as a consequence of enhanced oil recovery (±30%)



# Wellington, Kansas Demonstration

•Dolomite (Ca,Mg)CO<sub>3</sub>



## **Dissolution yields preferential flow paths in more heterogeneous carbonate rocks**



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**Figure 12:** Qualitative correlation between permeability contrast ( $k_f/k_i$ , increasing towards the right) and evolution of dissolution patterns from stable to less stable.

# Model parameters are constrained by characterization, pressure, and solution data



## **Reactive Transport Model**

#### Reactions

calcite + H<sup>+</sup> = Ca<sup>++</sup> + HCO<sub>3</sub><sup>-</sup> dolomite + 2H<sup>+</sup> = Ca<sup>++</sup> + Mg<sup>++</sup> + 2HCO<sub>3</sub><sup>-</sup>  $CO_{2(aq)} + H_2O = H^+ + HCO_3^-$ MgHCO<sub>3</sub><sup>+</sup> = Mg<sup>++</sup> + HCO<sub>3</sub><sup>-</sup> CaCO<sub>3(aq)</sub> + H<sup>+</sup> = Ca<sup>++</sup> + HCO<sub>3</sub><sup>-</sup> CaHCO<sub>3</sub><sup>+</sup> = Ca<sup>++</sup> + HCO<sub>3</sub><sup>-</sup>

#### **Mineral Reaction Rates**

$$\frac{dn}{dt} = -Sk_{298.15K}e^{-\frac{E}{R}\left(\frac{1}{T} - \frac{1}{298.15}\right)} \left(1 - \frac{Q}{K}\right)$$

Permeability-Porosity n – best fit

$$K_t = K_0 \left(\frac{\phi_t}{\phi_0}\right)^n$$

Surface Area-Porosity m - 2/3

$$S_{t} = S_{0} \left( \frac{\theta_{t}}{\theta_{0}} \frac{\phi_{t}}{\phi_{0}} \right)^{m}$$

## **Evolution of permeability is tied to the heterogeneity and the mineral reactivity**



## Mineral dissolution rates vary by 100 times and may require calibration of reactive surface area



$$\frac{dn}{dt} = -Sk_{298.15K}e^{-\frac{E}{R}\left(\frac{1}{T} - \frac{1}{298.15}\right)} \left(1 - \frac{Q}{K}\right)$$

## Validation Study – Big Sky Demonstration, Duperow Formation (Lee Spangler and Stacey Fairweather)



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www.bigskyco2.org/kevin\_dome\_site\_characterization

# How do you scale lab experiments to the field?

# Larger grid size reduces the permeability change



bulk permeability increase

# Calibration of down hole logs to better estimate variable permeability with depth in carbonate reservoirs





# NMR signal can be used to estimate down hole permeability

## Weyburn,

Wellington, Kansas



# Predicted permeability differs by orders of magnitude using standard value of *A*



**Calibrate** 
$$A = \frac{\rho^2}{\varphi^3 v \tau^2}$$
 **using independent**  
**measures**

- φ : porosity (Nuclear Magnetic Resonance)
- v : pore shape factor (2.5 for elliptical pores)
- τ : tortuosity (X-Ray Tomography, Nuclear Magnetic Resonance)
- ρ: surface relaxivity (Calibrated Nuclear Magnetic Resonance)

Daigle and Dugan JGR 2011

# **Tortuosity** (*τ*) is extracted from high resolution tomography images and the NMR porosity



$$A = \frac{\rho^2}{\varphi^3 v \tau^2}$$

- Matrix porosity assessed by difference between XRCT and NMR porosity
- Use a random walk algorithm to extract tortuosity from segmented pore network

Test – Initial estimates of caprock-like permeability from SDR equation and standard A is due to high Fe concentrations

- Solve for A = 5.33 x 10<sup>-09</sup> m<sup>2</sup>/s<sup>2</sup>
- NMR Porosity;  $\phi = 21.7\%$

 $\mathbf{k} = \mathbf{A} \cdot \mathbf{T}_{2,\text{LM}}^2 \cdot \boldsymbol{\varphi}^4$ 

- NMR T<sub>2,LM</sub>
- Measured Permeability; k = 0.027 mD
- Solve for Relaxivity; ρ = 65.6 μm/s
  - Standard for carbonates i
  - Reflects high paramagnet conten
- NMR Porosity;  $\varphi = 21.6\%$
- XRCT Tortuosity;  $\tau = 3.53$  m/m
- Pore shape factor; υ= 2.5 m<sup>2</sup>/m<sup>2</sup>
  - elliptical pores
  - could be refined with XRCT data

Weyburn flow units  

$$0.1$$
  
 $0.01$   
 $0.001$   
 $0.0001$   
 $0.0001$   
 $0.0001$   
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 $0.01$   
 $0.1$   
 $1$   
 $10$   
Log Mean T<sub>2</sub> (ms)

$$A = \frac{\rho^2}{\varphi^3 \upsilon \tau^2}$$

# Surface relaxivity (ρ) depends on mineralogy and Mn and Fe content



But p cannot resolve difference between estimated and measured permeability



# **Next steps in the calibration**



$$\mathbf{k} = \mathbf{A} \cdot \mathbf{T}_{2,\text{LM}}^2 \cdot \boldsymbol{\varphi}^4$$

$$A = \frac{\rho^2}{\varphi^3 \upsilon \tau^2}$$

- Measure the Fe/Mn content for all samples
- Conduct a sensitivity study of the parameters and power functions



#### Schlumberger, "Carbonate Reservoirs," 2007.

# Synergy

- Weyburn-Midale Carbon Storage Demonstration
- Wellington, Kansas Carbon Storage Demonstration
- Big Sky Carbon Storage Demonstration



# **Summary and Future Plans**

- Derived key reactive-transport parameters and their ranges for carbonate rocks over a wide range of heterogeneity and initial permeability
- Conducting a validation study using core from an independent CO<sub>2</sub> storage formation
- Developing a protocol for calibrating the NMR signal to provide meaningful in-situ permeability measurements
- Using numerical methods to scale laboratory parameters to reservoir
- Write final topical report on CO<sub>2</sub> storage potential in carbonate rocks.



# **Bibliography**

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