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Fluid-Rock Interactions at the Interface between Reservoir Rock and Cap Rock: An Experimental Case Study of Mineral Trapping

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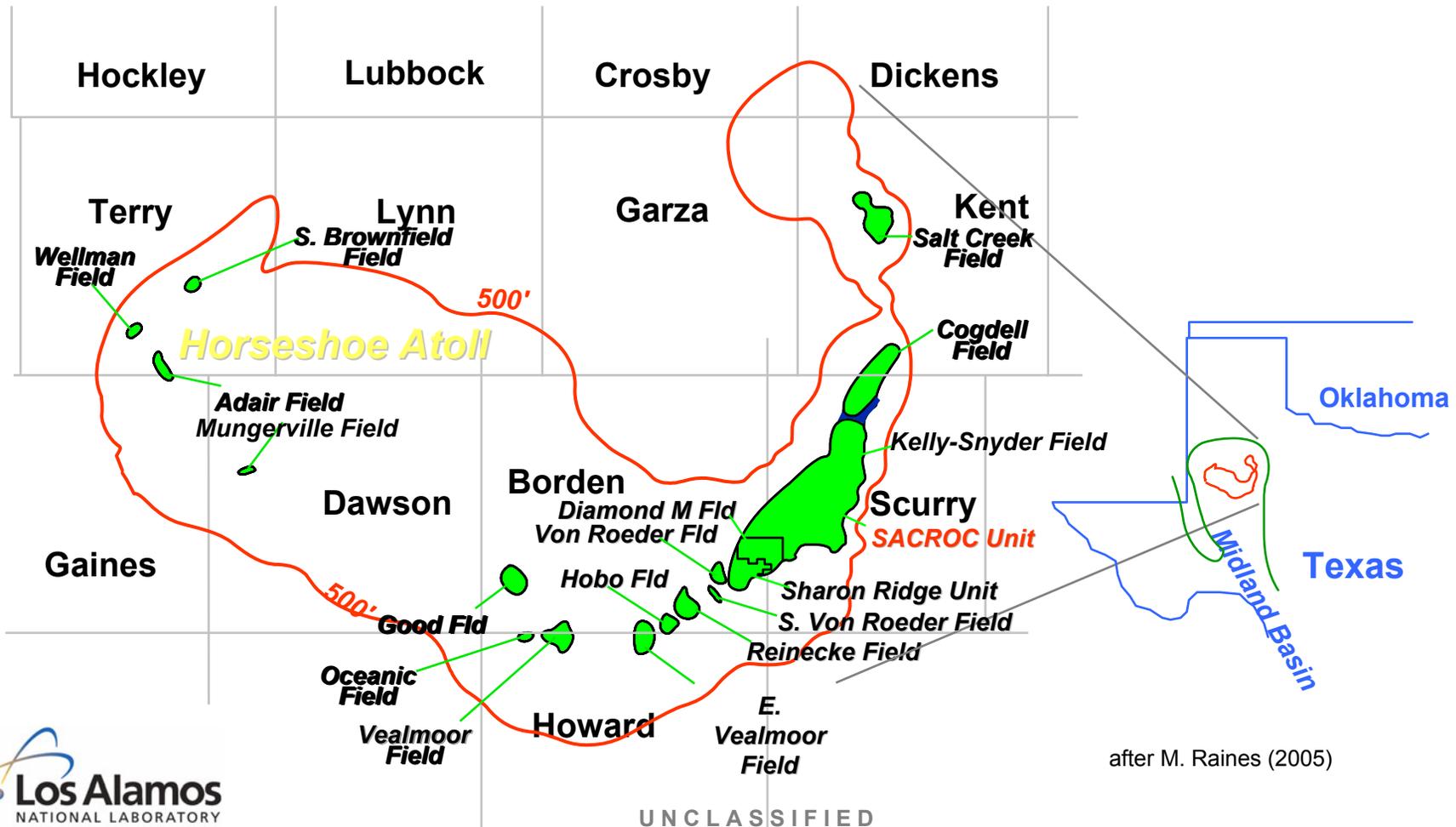
OUTLINE

- Introduction
- Core flood experiments
- Modeling and interpretation

GEOCHEMICAL TRAPPING

- Solubility trapping
 - $\text{CO}_2(\text{gaseous}) + \text{H}_2\text{O} \rightarrow \text{H}_2\text{CO}_3(\text{aq})$
- Ionic trapping
 - $3 \text{ K-spar} + 2\text{H}_2\text{O} + 2\text{CO}_2 \rightarrow \text{Muscovite} + 6 \text{ Quartz} + 2\text{K}^+ + 2\text{HCO}_3^-$
 - $\text{H}_2\text{CO}_3(\text{aq}) + \text{OH}^-(\text{aq}) \rightarrow \text{HCO}_3^-(\text{aq})$
 - $\text{HCO}_3^-(\text{aq}) + \text{OH}^-(\text{aq}) \rightarrow \text{CO}_3^{2-}(\text{aq})$
- Mineral Trapping
 - $\text{CO}_3^{2-}(\text{aq}) + \text{Ca}^{++} \rightarrow \text{CaCO}_3(\text{solid})$
 - $\text{CO}_3^{2-}(\text{aq}) + \text{Fe}^{++} \rightarrow \text{FeCO}_3(\text{solid})$
 - $\text{CO}_3^{2-}(\text{aq}) + \text{Mg}^{++} \rightarrow \text{MgCO}_3(\text{solid})$
 - $\text{CO}_2(\text{aq}) + \text{K-spar} + \text{Na}^+ \rightarrow \text{NaAl}(\text{CO}_3)(\text{OH})_2(\text{Solid}) + \text{SiO}_2 + \text{K}^+$

SACROC LOCATION MAP



after M. Raines (2005)

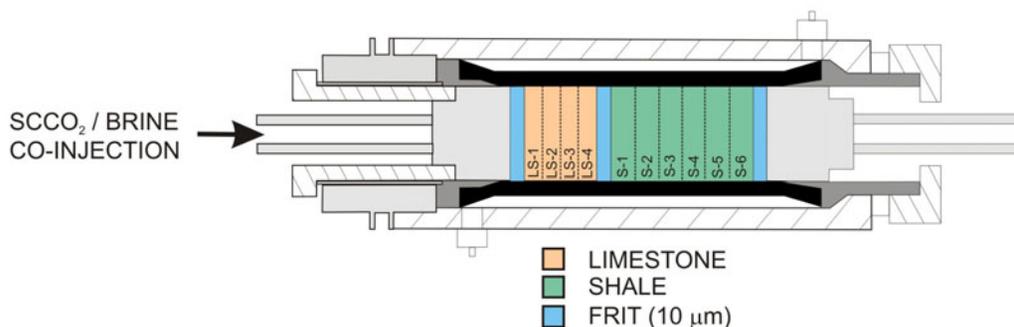


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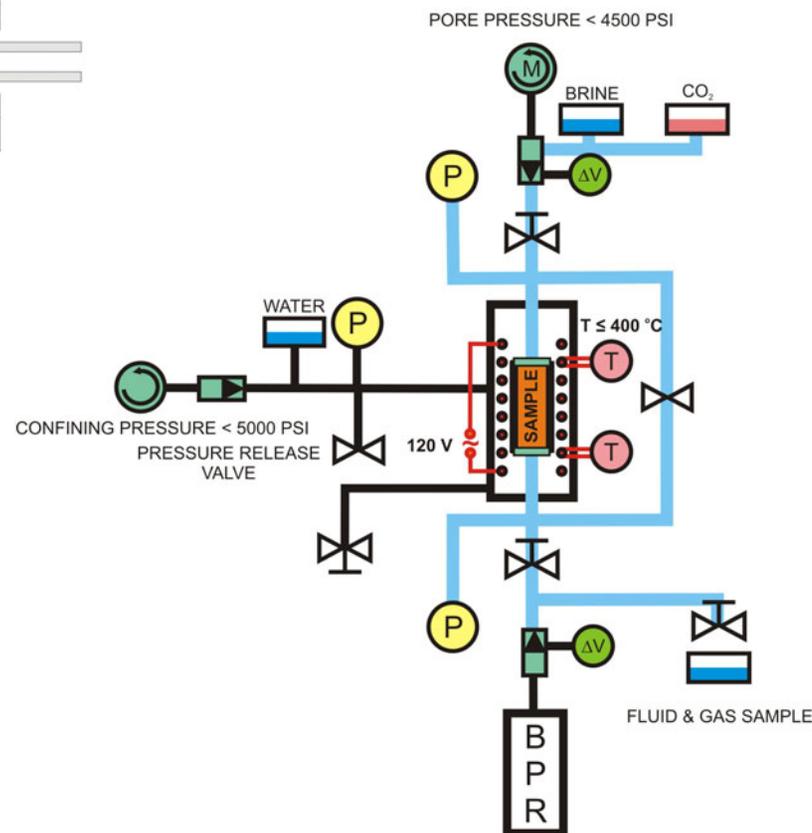
Modified from Vest (1970)



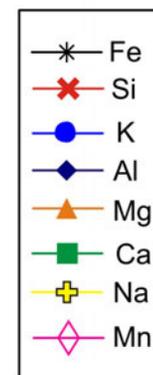
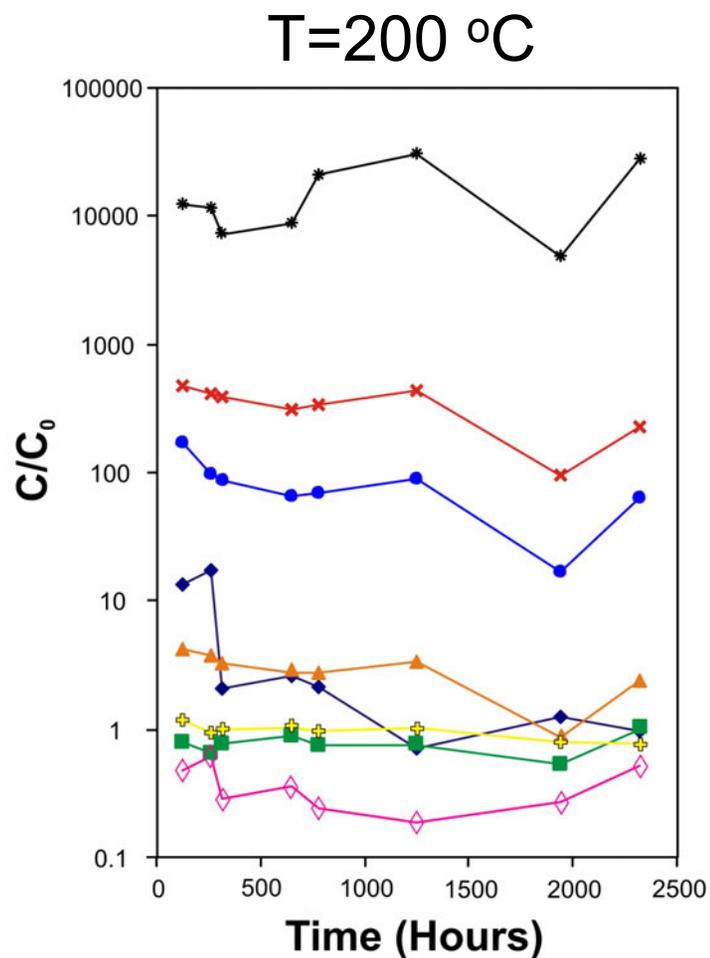
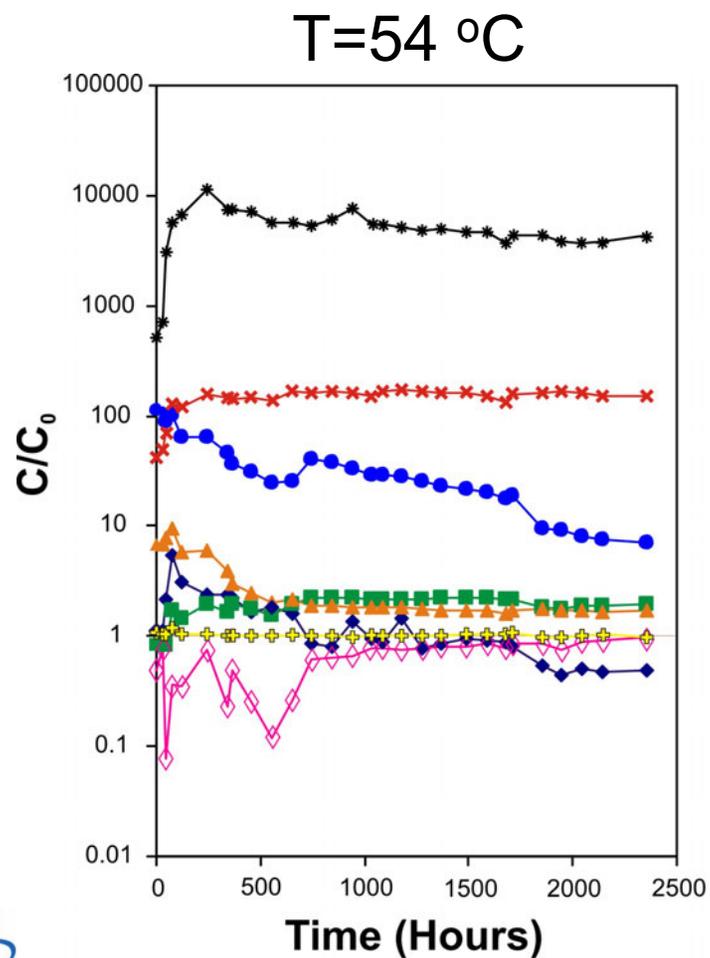
EXPERIMENTAL SETUP



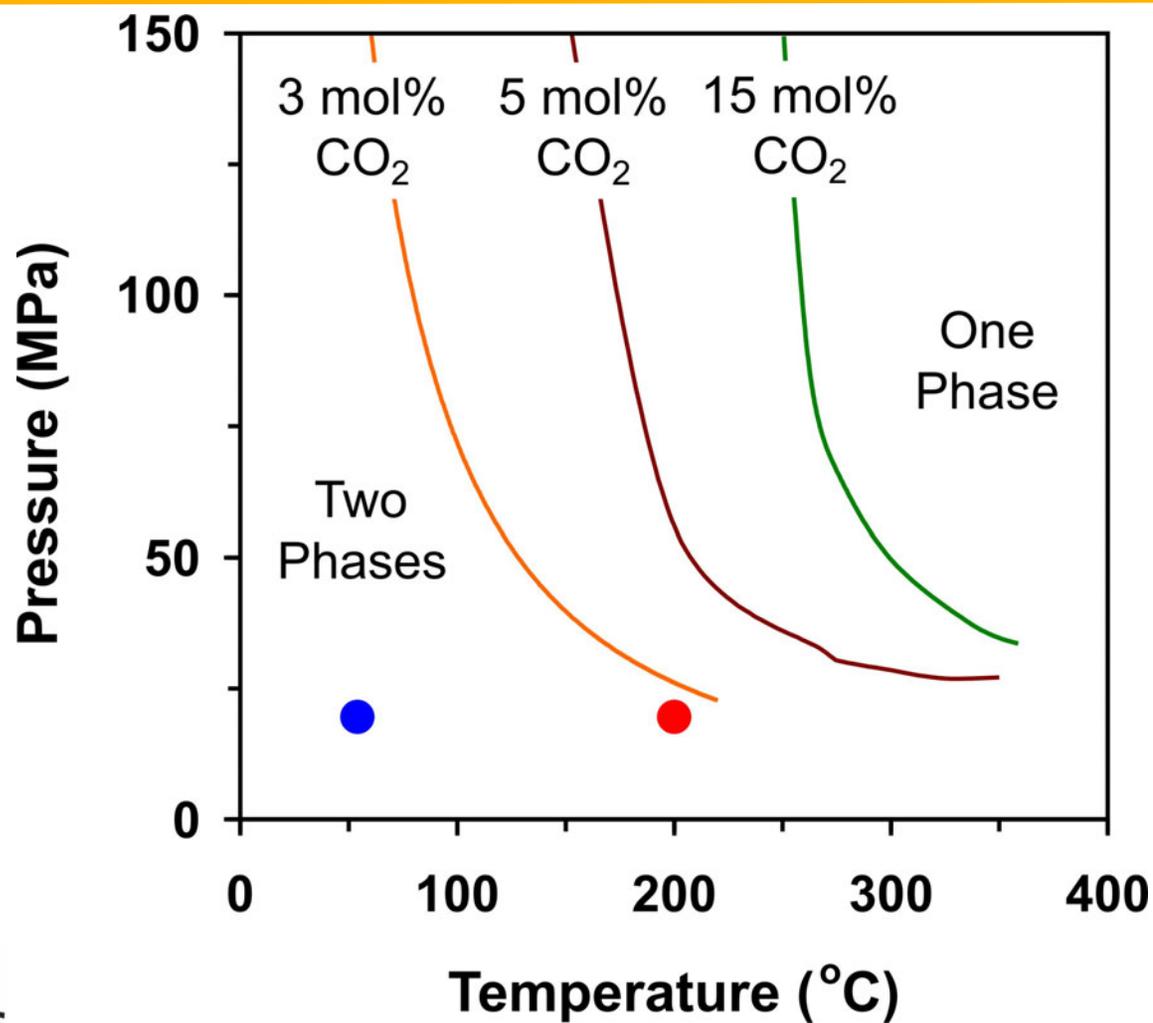
	EXPERIMENT I	EXPERIMENT II
TEMPERATURE	54 °C	200 °C
PRESSURE	2880 PSI	2880 PSI
DURATION	3263 HOURS	3263 HOURS
FLUIDS	SCCO ₂ & BRINE	SCCO ₂ & BRINE
ROCKS	LIMESTONE ILLITE-RICH SHALE	LIMESTONE ILLITE-RICH SHALE
FLOWRATE	SCCO ₂ : 4 μl min ⁻¹ BRINE: 8 μl min ⁻¹	SCCO ₂ : 4 μl min ⁻¹ BRINE: 8 μl min ⁻¹



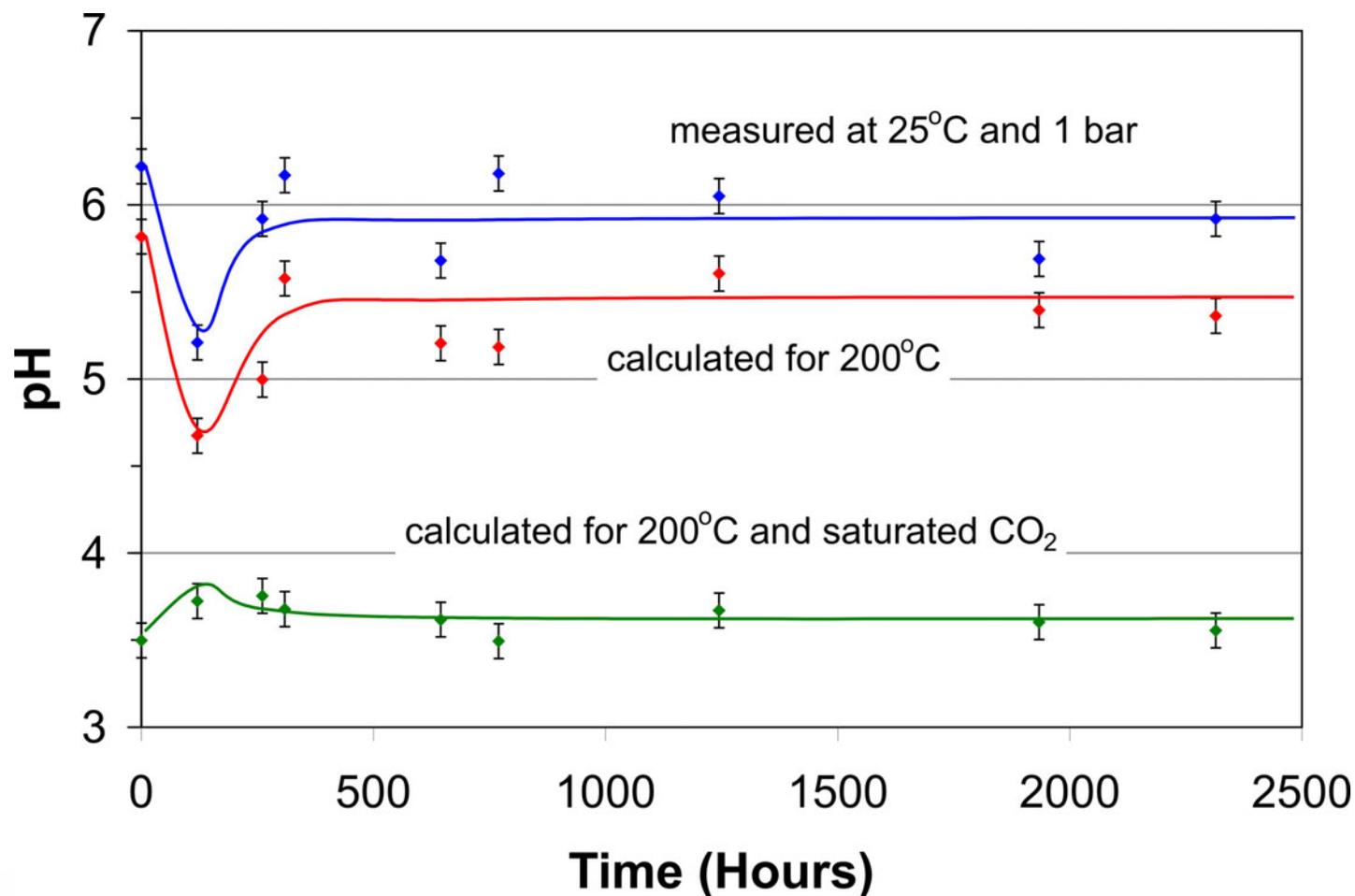
FLUID CHEMISTRY



H₂O-CO₂ PHASE SPACE



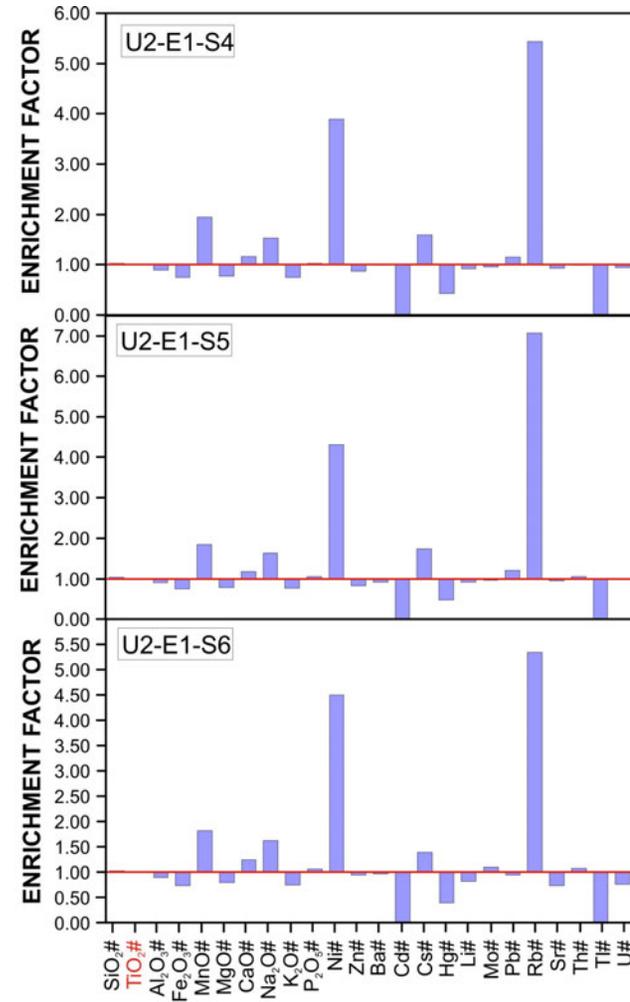
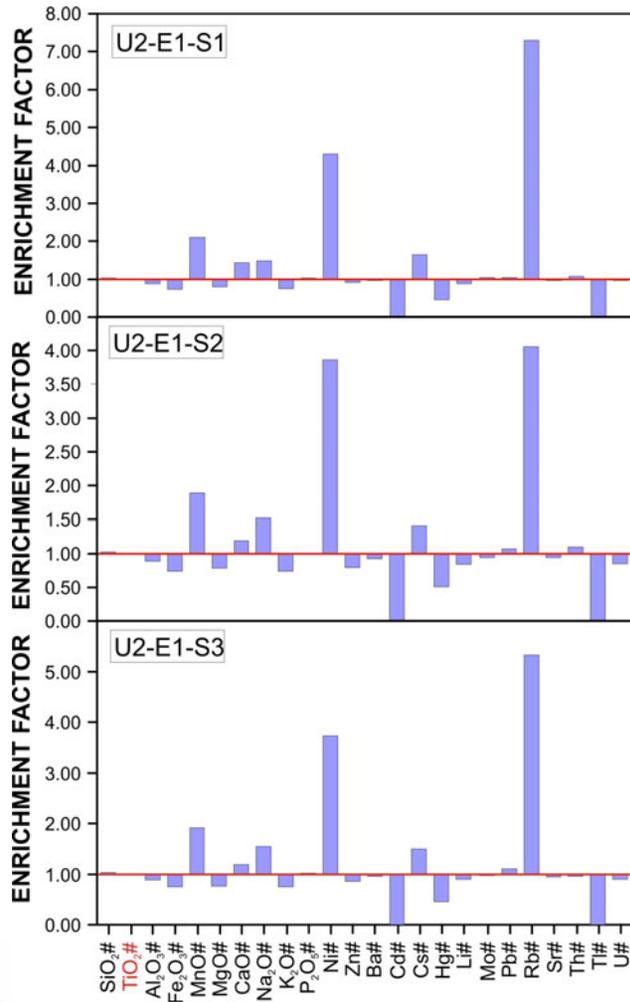
pH EVOLUTION OF BRINE (200°C EXPERIMENT)



CHEMICAL CHANGES – ENRICHMENT FACTOR

54 °C EXPERIMENT

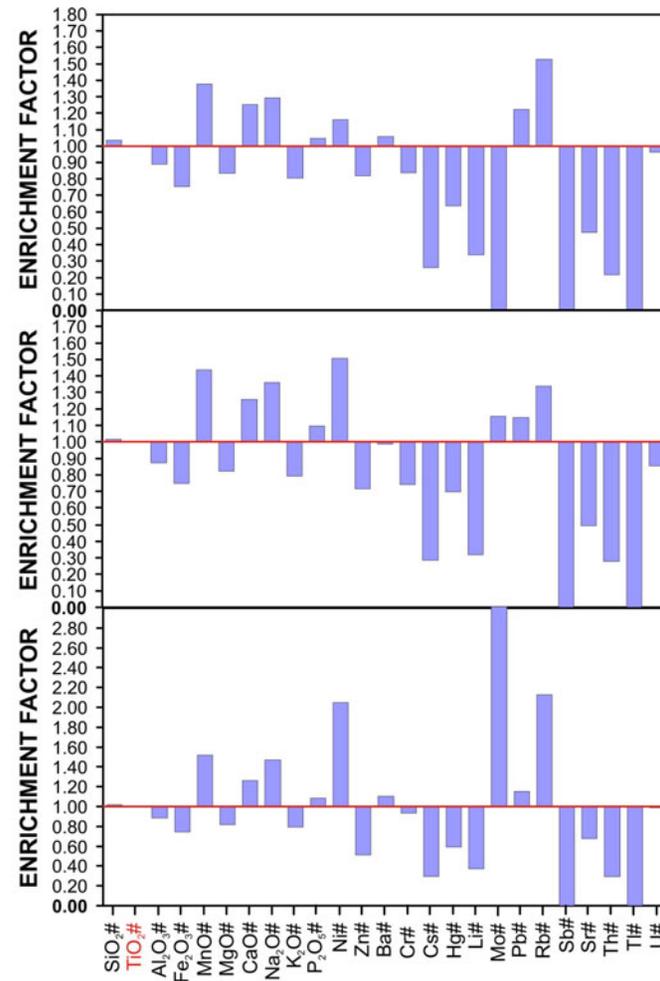
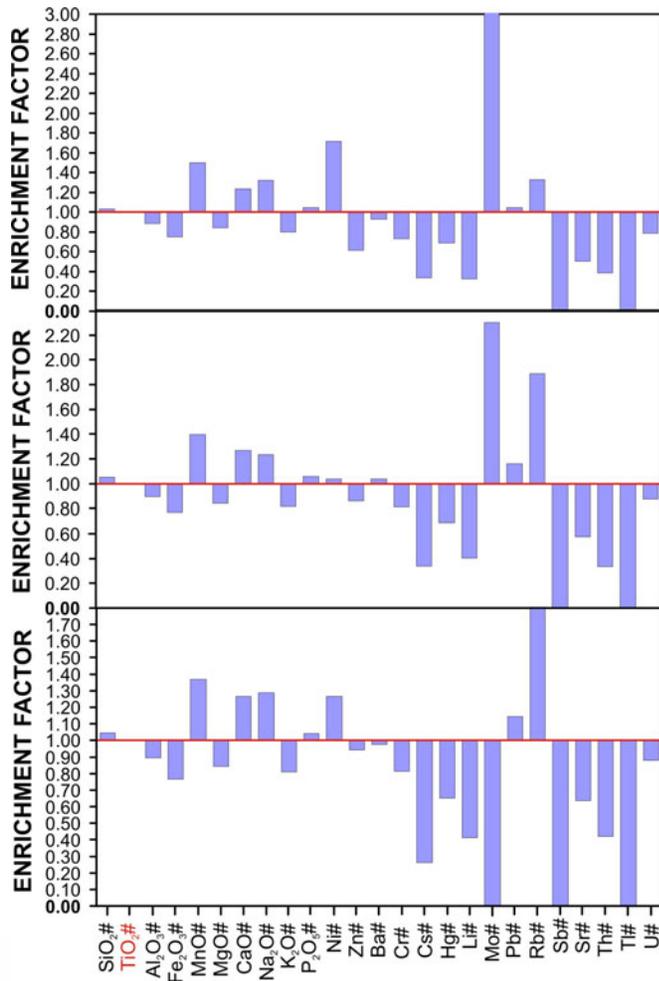
$$X = [A(X) / A(\text{TiO}_2)] / [L(X) / L(\text{TiO}_2)]$$



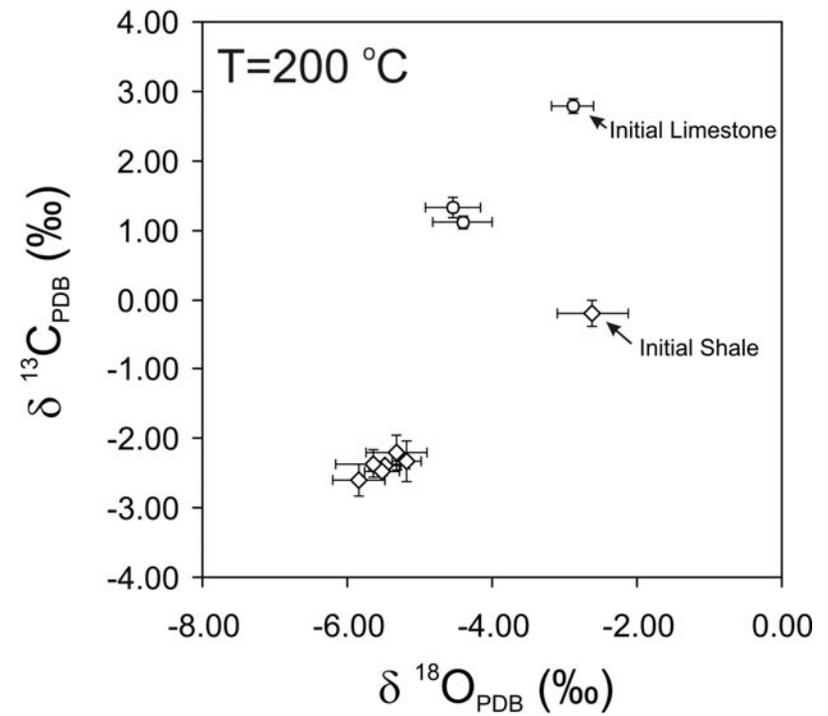
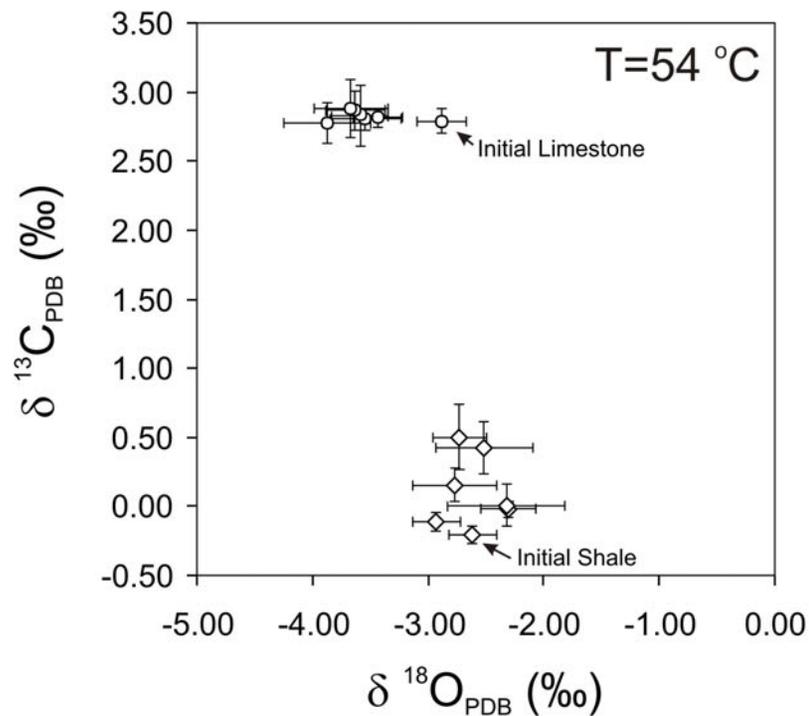
CHEMICAL CHANGES – ENRICHMENT FACTOR

200 °C EXPERIMENT

$$X = [A(X) / A(\text{TiO}_2)] / [L(X) / L(\text{TiO}_2)]$$

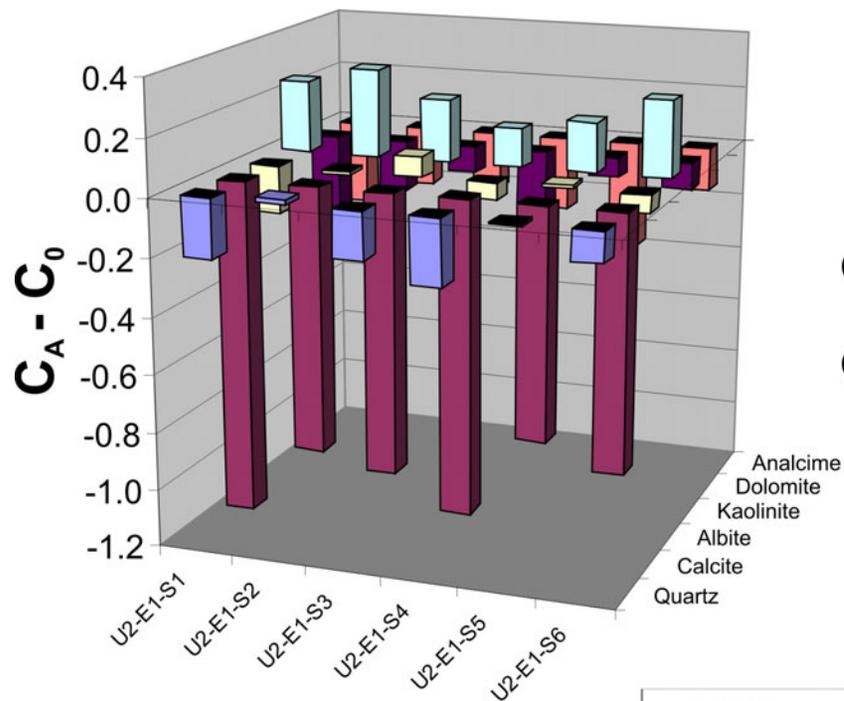


STABLE ISOTOPES

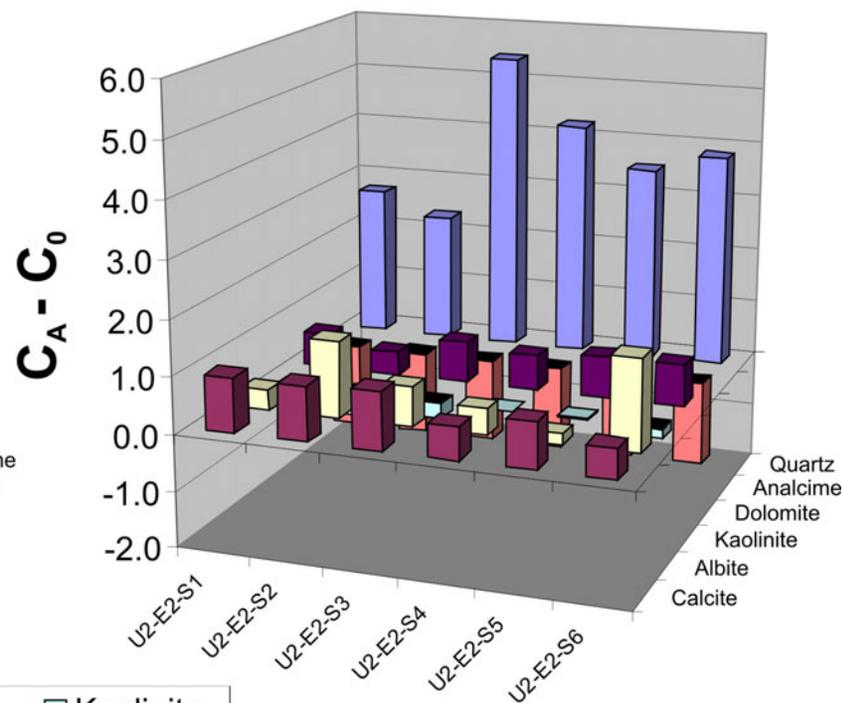


MINERALOGY

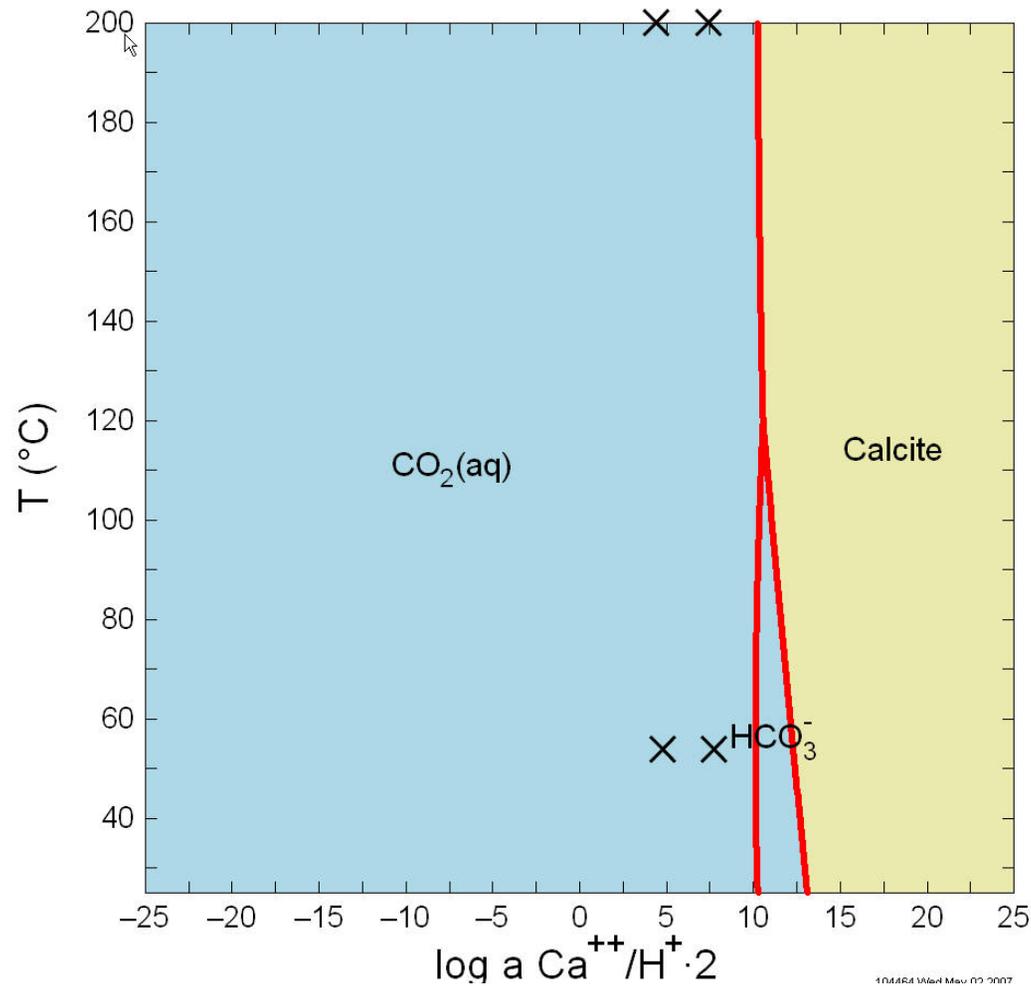
T = 54 °C



T = 200 °C



GEOCHEMICAL MODELING



CONCLUSIONS

- The experiment performed at in-situ reservoir conditions (54 °C / 2880 psi) was dominated by the dissolution of rock-forming minerals such as calcite whereas calcite and quartz precipitated at elevated temperatures (200 °C / 2880 psi).
- The isotopic signature of $\delta^{13}\text{C}$ is not affected by the mechanism controlling the dissolution of carbonates at 54 °C but shows a strong change due to precipitation of calcite during the experiment performed at 200 °C.
- Carbon and oxygen isotopes have been shown to be sensitive to small amounts of carbonate precipitates.

ACKNOWLEDGEMENTS

SPECIAL THANKS TO:

Toti Larson

Melissa Fittipaldo

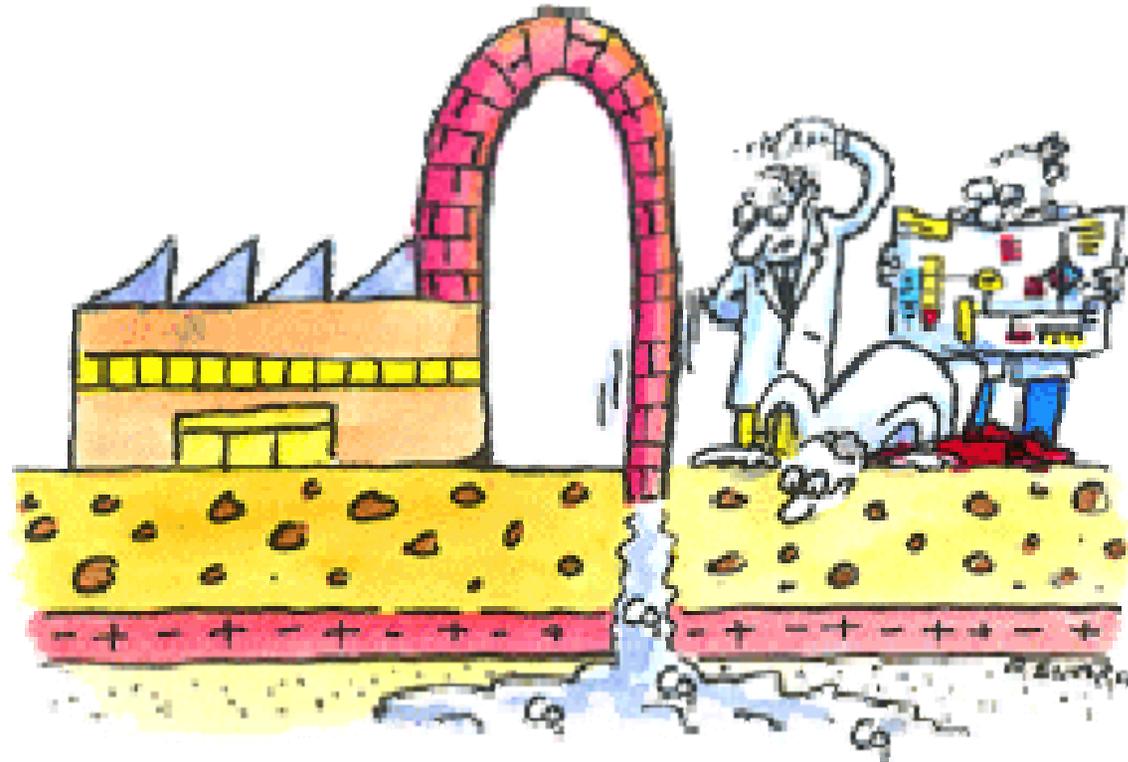
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Emily Kluk



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Thank you very much for your attention!

